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Bayesian geostatistical modelling of high-resolution *NO*₂ exposure in Europe combining data from monitors, satellites and chemical transport models



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ABSTRACT

Bayesian geostatistical regression (GR) models estimate air pollution exposure at high spatial resolution, quantify the prediction uncertainty and provide probabilistic inference on the exceedance of air quality thresholds. However, due to high computational burden, previous GR models have provided gridded ambient nitrogen dioxide (NO_2) concentrations at smaller areas of investigation. Here, we applied these models to estimate yearly averaged NO₂ concentrations at 1 km² spatial resolution across 44 European countries, integrating information from in situ monitoring stations, satellites and chemical transport model (CTM) simulations. The tropospheric values of NO2 derived from the ozone monitoring instrument (OMI) onboard the National Aeronautics and Space Administration's (NASA's) Aura satellite were converted to near ground NO_2 concentration proxies using simulations from the 3-D global CTM (GEOS-Chem) at $0.5^{\circ} \times 0.625^{\circ}$ spatial resolution and surface-to-column NO2 ratios. Simulations from the Ensemble of regional CTMs at spatial resolution of $0.1^{\circ} \times 0.1^{\circ}$ were extracted from the Copernicus atmosphere monitoring service (CAMS). The contribution of these covariates to the predictive capability of geostatistical models was for the first time evaluated here through a rigorous model selection procedure along with additional continental high-resolution satellite-derived products, including novel data from the pan-European Copernicus land monitoring service (CLMS). The results have shown that the conversion of columnar NO₂ values to surface quasi-observations yielded models with slightly better predictive ability and lower uncertainty. Nonetheless, the use of higher resolution CAMS-Ensemble simulations as covariates in GR models granted the most accurate surface NO₂ estimates, showing that, in 2016, 16.17 (95% C.I. 6.34–29.96) million people in Europe, representing 2.97% (95% C.I. 1.16% - 5.50%) of the total population, were exposed to levels above the EU directive and WHO air quality guidelines threshold for NO_2 . Our estimates are readily available to policy makers and scientists assessing the burden of disease attributable to NO2 in 2016.

1. Introduction

Ground-level nitrogen dioxide (NO_2) concentrations represent a serious public health concern. Exposure to elevated levels of NO_2 are associated with increased cardiovascular, respiratory and all-cause mortality and respiratory morbidity (WHO, 2013; European Environment Agency, 2014). Many early epidemiological studies used measurements from the nearest monitoring sites to estimate NO_2 exposure (Hesterberg et al., 2009; Latza et al., 2009). This approach introduces exposure misclassification (Jerrett et al., 2005), since it fails to capture the sources of local spatial variability. Air quality related policies require an accurate assessment of the exposures in any given area. In practice, this is difficult to address due to the complexity of the processes involved and due to the sparsity of the monitoring data. In Europe, air quality monitoring is maintained by the European Environment Agency's (EEA's) member states, and although there is a relatively dense network of stations, when compared to other areas globally, large parts of the continent remain unmonitored. Therefore, predicting the spatial distribution of the outdoor air pollution is an important research goal for environmental health.

There are a number of methods that have been used to provide gridded NO_2 estimates on large scales (like continental or global). These include chemical transport models (CTMs) and empirical models based on their outputs (Schaap et al., 2008; Lamsal et al., 2008; Geddes et al., 2015), spatial interpolation, such as kriging (Beelen et al., 2009; Young et al., 2016) and land-use regression (LUR) (Beelen et al., 2009; Novotny et al., 2011; Vienneau et al., 2013; Knibbs et al., 2014; Bechle et al., 2015; De Hoogh et al., 2016; Young et al., 2016; Larkin et al.,

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2017). Each of these models has both strengths and weaknesses. In particular, the continental or global-scale dispersion models (e.g. CTMs), which follow the principles of chemistry and physics, usually have low spatial resolution, due to limitations related to computational capacity and coarse resolution of the emission inventories. The latter are typically available at 4° \times 5°, 2.0° \times 2.5°, 0.500° \times 0.667° or $0.2500^{\circ} \times 0.3125^{\circ}$ scales (Di et al., 2017); however, higher resolution inventories, like EDGAR (Emission Database for Global Atmospheric Research - https://edgar.jrc.ec.europa.eu/) are also becoming available in recent time (Crippa et al., 2018). The LUR models are technically less challenging and computationally easy to fit. They can provide predictions of NO₂ concentration at very high spatial resolution (Vienneau et al., 2013), however, they don't take into account the spatial correlation present in the air pollution data, thus overestimating the significance of the covariates. Moreover, their predictive ability is lower when compared to more complex geostatistical or geographically weighted regression models (Beloconi et al., 2018). Bayesian geostatistical regression (GR) models capture the spatial correlation present in the pollutant concentrations and provide estimates of the prediction uncertainty. Furthermore, they allow a straightforward assessment of the exposure burden through high-resolution population estimates, since the posterior predictive distributions of NO2 can be derived for every pixel within the study domain. However, as for any statistical model, predictions rely on appropriate input data (i.e. monitoring stations), which is available for most of Europe but may not be available for other regions of the world. In addition, as discussed in Shaddick et al. (2013), the prediction of NO_2 levels using this set of models at high spatial resolution over large scales (e.g. for 15 countries of the European Union (EU)) is computationally complex. The approximate Bayesian inference using the stochastic partial differential equations (SPDE) approach and integrated nested Laplace approximation (INLA) (Rue et al., 2009; Lindgren et al., 2011) has shown to provide yearlyaveraged gridded concentrations of particulate matter at 1 km² spatial resolution across Europe in a reasonable computational time (Beloconi et al., 2018).

Most of the data-driven air-quality assessments incorporate geographical covariates derived from satellite-based observations, since they usually provide spatial coverage over the entire domain of interest. Several previous works have used the tropospheric NO_2 from the ozone monitoring instrument (OMI), onboard the National Aeronautics and Space Administration's (NASA's) Aura satellite (Levelt et al., 2006), to estimate the corresponding surface concentrations (Lamsal et al., 2008; Novotny et al., 2011). However, there are two main issues that should be taken into consideration while using this data for modelling.

First, the satellite instruments measure total tropospheric columns and therefore the NO₂ proxies derived from OMI may not represent well the corresponding surface concentrations. Aerial measurements reveal that the concentration of NO₂ in the tropospheric column is determined primarily by NO_2 in the mixed layer, as well as by NO_2 in the boundary layer (Martin et al., 2004; Richter et al., 2005; Martin et al., 2006; Boersma et al., 2008; Bucsela et al., 2008). However, the proportion of NO_2 in these two layers varies in both space and time (Lamsal et al., 2008; Grajales and Baquero-Bernal, 2014). Although recent works suggested that the satellite column abundance (total concentration within a vertical column) may be efficient to track spatial patterns in the ground-level NO₂ (Bechle et al., 2013; Knibbs et al., 2014; Bechle et al., 2015), in which case conversion of columnar values to surface concentrations becomes unnecessary within a LUR model, the contribution of this conversion within a GR modelling framework to the best of our knowledge has not been evaluated.

Second, although OMI's spatial resolution (of up to 13×24 km at nadir) is the highest among all available space-borne NO_2 sensors (except of the recently released data from the Sentinel 5 Precursor tropospheric monitoring instrument (S5P/TROPOMI) which provides measurements at 7 × 3.5 km² resolution (Veefkind et al., 2012) from late 2017 onwards), it is still too coarse to capture near-source (e.g. near-

roadway) concentration variability (Novotny et al., 2011). Since many parameters, including local combustion sources, land surface characteristics and atmospheric conditions influence NO_2 formation and dispersion (Bechle et al., 2015; Young et al., 2016; Larkin et al., 2017), the use of statistical models which incorporate covariates of higher spatial-resolution could allow estimation of the local-scale variation of the pollutant. Additionally, in the view of the new S5P/TROPOMI data, it is of interest to evaluate whether higher resolution CTM simulations will lead to even better predictive ability of the statistical models, when considered as additional covariates.

Here, we applied Bayesian GR models to estimate NO2 gridded concentrations at 1 km² resolution across 44 European countries. For the first time we evaluated the contribution of the tropospheric columnto-suface conversion of satellited-based NO2 proxies within a geostatistical modelling framework. We incorporated the vertical distribution of NO2 derived from the global 3-D CTM (GEOS-Chem, Bey et al., 2001), to quantify the corresponding tropospheric columns and to infer quasi-observed surface concentrations from the columnar OMI measurements using the estimated surface-to-column NO2 ratios. Additionally, we included a large set of high-resolution geo-referenced predictors available at continental scale, such as the novel pan-European Copernicus land monitoring products (CLMS, 2019) and meteorological data and tested the contribution of each predictor through a rigorous model selection procedure. We compared the results with simulations from the Ensemble of regional CTMs available at the Copernicus atmosphere monitoring service (CAMS, 2019) at higher spatial resolution than the OMI/GEOS-Chem estimates. The Bayesian framework allowed us to quantify the uncertainty of the predictions and to determine areas that exceed the air quality guidelines (AQGs) thresholds set by the European Union (EU) and World Health Organization (WHO), as well as to assess the number of people living in such areas. This study provides relevant information for policy and decision-makers in Europe and can contribute to improved estimates of the burden of disease attributable to NO₂ (Lim et al., 2012; Forouzanfar et al., 2016).

2. Materials and methods

2.1. Study area and data

The raw NO2 measurements were obtained from the Air Quality e-Reporting database (Air Quality e-Reporting, 2019) maintained through the European environment information and observation network (Eionet). The monitoring network covers up to 38 European Countries, including all EU member states and EEA cooperating countries. The database consists of multi-annual time series of air quality measurements for a number of pollutants together with the meta-information on the monitoring stations involved. Here, the analysis was based on the yearly averaged 24 h data of 2016 at sites with ≥75% data capture. The station data were used for both, model building and model validation. Fig. 1 illustrates the locations of the monitoring sites used in this work together with the yearly averaged measured concentrations of NO2. All data considered in this study were converted to the Lambert Azimuthal Equal Area (ETRS89-LAEA5210) projection recommended by the EEA (European Environment Agency, 2006) for storing raster data, statistical analysis and mapping purposes.

The satellite-derived product of columnar NO_2 derived from OMI offers near global daily coverage of NO_2 column abundance at spatial resolution of up to 13 × 24 km at nadir. The level-3 daily data of tropospheric NO_2 concentration (OMNO2d) (Krotkov, 2013), cloud-screened at 30% and binned into $0.25^{\circ} \times 0.25^{\circ}$ grids, were accessed from the NASA's Goddard Earth Sciences Data and Information Services Center (GES DISC) website (GES DISC, 2019). To obtain pollutant concentrations at the surface, the surface-to-column NO_2 ratios (Lamsal et al., 2008, 2010; Bechle et al., 2011) derived from the global 3-D atmospheric model (GEOS-Chem v.11-01, http://acmg.seas.harvard.



Fig. 1. Nitrogen Dioxide. Annual average concentration of NO₂ in 2016 at 2852 monitoring sites across Europe.

edu/geos/) were used (see Section 2.2 for details).

Satellite-based *NO*₂ estimates reflect contributions from all sources, including e.g. emissions from industry, roads, airports or harbours and are not explicitly included as covariates within statistical models (Novotny et al., 2011). Following our previous work (Beloconi et al., 2018), a number of additional products were used as predictors in our models, to better assess the spatial variability of pollutant's

concentrations across Europe. The choice of variables was guided by literature review and data availability on continental scale. Table 1 summarizes the covariates used in this work. Detailed information on data sources, retrieval and processing is given in the appendix.

Table 1

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Product	Temporal Resolution	Spatial Resolution	Source
Tropospheric cloud-screened NO_2 (Ω_O)	daily (13:00-15:00 LT)	$0.25^{\circ} imes 0.25^{\circ}$	OMI
Surface NO_2 estimates (S_0)	daily (13:00–15:00 LT)	$0.5^{\circ} imes 0.625^{\circ}$	OMI/GEOS-Chem
Surface NO_2 estimates (S_{ENS})	year 2016	$0.1^{\circ} imes 0.1^{\circ}$	CAMS - Ensemble
Corine Land Cover (LC)	year 2012	100 m × 100 m	CLMS
Tree Cover Density (TCD)	year 2015	20 m × 20 m	CLMS
Imperviousness (IMP)	year 2015	20 m × 20 m	CLMS
Digital Elevation Model (DEM)	year 2000	30 m × 30 m	EEA
Night Time Lights (NTL)	year 2013	$1 \text{ km} \times 1 \text{ km}$	NOAA
Land Surface Temperature Day & Night (LST)	2 acquisitions per day	$1 \text{ km} \times 1 \text{ km}$	MODIS Aqua and Terra
Normalized Difference Vegetation Index (NDVI)	2 acquisitions per day	$1 \text{ km} \times 1 \text{ km}$	MODIS Aqua and Terra
Road Density (RD)	February 2016	$1 \text{ km} \times 1 \text{ km}$	OpenStreet Maps
Specific Humidity (SHUM)	every 6 h	$0.2^{\circ} imes 0.2^{\circ}$	NCEP/CFSv2
Precipitation (PREC)	every 6 h	$0.2^{\circ} imes 0.2^{\circ}$	NCEP/CFSv2
Wind Speed (WS)	every 6 h	$0.2^{\circ} imes 0.2^{\circ}$	NCEP/CFSv2
Distance to Sea (DISS)	year 2015	vector	EEA
Distance to Roads (DISR)	February 2016	vector	OpenStreet Maps

2.2. Chemical transport models

Conversion of the NO2 proxies derived from OMI to surface concentrations require information about the NO2 vertical profile in the troposphere. Here we used the nested-grid version of the GEOS-Chem (Wang et al., 2004; Zhang et al., 2012) model which is operated at $0.5^{\circ} \times 0.625^{\circ}$ spatial resolution with 47 vertical layers and a transport and chemistry time step of 10 and 20 min, respectively. The boundary conditions are updated every 3 h using global simulations from a $2^{\circ} \times 2.5^{\circ}$ parent model (one-way nesting). We used assimilated meteorological fields (Modern-Era Retrospective Analysis for Research and Application (MERRA2)) provided by the NASA's Global Modelling and Assimilation Office. One year spin-up for 2015 was performed to produce simulations for 2016. Daily tropospheric NO2 columns, using 2 h average between 13:00 and 15:00 local time (LT), corresponding to the time of the Aura satellite overpass, were generated to compare with the OMI retrievals. In particular, the simulated NO₂ mixing ratios obtained from GEOS-Chem at 47 vertical layers were integrated from the surface (i.e. the lowest layer of the model, which is approximately 60 m above ground) to the modelled tropopause for every pixel and day within the study domain. Ideally for CTMs the inner product of the simulated profiles and the OMI averaging kernels should be taken, before comparing the modeled retrieval equivalents to the OMNO2d product. When the model profile shape is different from the a priori profile used in the satellite retrieval, this comparison could have an impact on the agreement between the estimates using these two approaches (Eskes et al., 2003). However, when averaged over larger regions the averaging-kernel based columns are very similar to the direct columns, showing only a remarkably small bias when not taken into account (Huijnen et al., 2010). Therefore, here the OMI product is directly compared to the modelled total columns.

Additionally, CTM simulations of surface NO₂ concentration were obtained from the regional production of the Copernicus atmosphere monitoring service (CAMS) reanalysis dataset of atmospheric composition produced by the European centre for medium-range weather forecasts (ECMWF) (Inness et al., 2019). The CAMS regional air quality production (CAMS, 2019) is based on the median value of 7 partner state-of-the-art numerical air quality models (Ensemble) operated by 8 European institutes: CHIMERE from INERIS (France), EMEP from MET Norway (Norway), EURAD-IM from University of Cologne (Germany), LOTOS-EUROS from KNMI and TNO (Netherlands), MATCH from SMHI (Sweden), MOCAGE from METEO-FRANCE (France) and SILAM from FMI (Finland). Common to all models are the meteorological parameters' settings (coming from the ECMWF global weather operating system), the boundary conditions for chemical species (coming from the CAMS IFS-MOZART global production) and the emissions coming from CAMS (for anthropogenic emissions over Europe and for biomass burning). The CAMS-Ensemble simulations used in this work represent annual averages of surface NO₂ concentration (in μ g/m³) at 0.1° × 0.1° spatial resolution for the year 2016, and are denoted as SENS in all the further analyses.

2.3. Vertical correction of OMI data

The GEOS-Chem vertical profiles were used to estimate the OMI concentrations at the surface employing the following formula (Lamsal et al., 2008):

$$S_0 = \frac{S_G}{\Omega_G} \times \Omega_0 \tag{1}$$

where *S* represents the superficial level concentration, measured in parts per billion volume (ppbv) and Ω - the tropospheric *NO*₂ column, measured in molecules/cm². The sub-indices *O* and *G* indicate OMI and GEOS-Chem, respectively. The OMI-derived surface concentrations *S*₀ (ppbv) represent the mixing ratio in the lowest layer of the model.

We generated daily estimates of surface NO2, based on the 2 h

average (between 13:00–15:00 LT) surface-to-column ratios. Subsequently, annual averages of NO_2 abundances were calculated for every pixel within the study domain. The raw NO_2 data obtained from the EEA as well as the AQGs in Europe are based on measurements expressed in mass density (μ g/m³) units, therefore, the S_O estimates in *ppbv* were converted to μ g/m³ using the following formula:

$$1\frac{\mu g}{m^3} = 1 \text{ ppbv} \times \frac{M(NO_2)P}{RT}$$
(2)

where *R* is the gas constant (8.3144 J mol⁻¹ K⁻¹)), $M(NO_2)$ is the molar mass of nitrogen dioxide (46.0055 g mol⁻¹), *P* and *T* are the grid level surface pressure (in mPa), and temperature (in K), respectively. The EU legislation (EU, 2008) specifies that, for gaseous pollutants the volume must be standardized at a temperature of 293 K (20 °C) and an atmospheric pressure of 101,3 kPa. In this case, the temperature refers to the bench temperature of the instrument and the pressure to the internal pressure in the measurement cell, not to ambient conditions (Sofen et al., 2016). The use of these values resulted in a conversion factor of 1.9125 µg/m³ (ppb⁻¹).

2.4. Geostatistical modelling

In the geostatistical framework spatial correlation is modelled by location-specific random effects through a Gaussian process. The covariance matrix of this process assumes a correlation decay which is a function of distance between locations. Let Y_s represents the log-observed annual average of NO_2 concentration (calculated using daily averaged station measurements) at site s ($s = 1, ..., S \in D \subseteq \mathbb{R}^2$). We assumed a stationary, isotropic geostatistical regression (GR) model,

$$Y_s = \beta_0 + X_s \beta + \omega_s + \epsilon_s \tag{3}$$

where β_0 is the intercept term, β the $k \times 1$ vector of regression coefficients associated with X_s , w_s the spatial random effect, and ϵ_s the random error assumed to be i.i.d. from a $N(0, \sigma_{\epsilon}^2)$. All the continuous covariates were standardized by subtracting the mean and dividing by the standard deviation (calculated using the yearly averaged measurements from all the monitoring stations). For the estimation of model parameters, the covariates were extracted at the locations of the stations, while for the prediction at unknown locations each covariate was aggregated within a fixed 1 km² grid using bilinear or nearest neighbour interpolation methods (for continuous and categorical data, respectively).

We assumed that the spatial random effect $w_s = (w_1, ..., w_S)^T$ arises from a multivariate normal distribution:

$$\boldsymbol{\omega} \sim \mathcal{N}(\mathbf{0}_{S}, \, \sigma_{\omega}^{2} \mathbf{R}_{\omega}) \tag{4}$$

with $\mathbf{0}_S$ a $S \times 1$ vector of zeros and σ_{ω}^2 is the spatial process variance. \mathbf{R}_{ω} is the $S \times S$ dense correlation matrix with elements $(\mathbf{R}_{\omega})_{ij} = C(||\mathbf{s}_i - \mathbf{s}_j||)$, where $C(\cdot)$ is the Matern function given by:

$$C(d_{ij}) = \frac{1}{\Gamma(\nu)2^{\nu-1}} (\kappa d_{ij})^{\nu} K_{\nu}(\kappa d_{ij})$$
(5)

where d_{ij} is the distance between stations *i* and *j*, κ is a scaling parameter, ν a smoothing parameter (fixed to 1 in our application) and K_{ν} is the modified Bessel function of second kind. This specification implies that the range *r* (the distance at which the spatial variance becomes less than 10%) is given by $r = \frac{\sqrt{8\nu}}{r}$.

The Bayesian model formulation is completed by specifying prior distributions for the parameters and the hyperparameters. We wanted the corresponding posterior distribution to be solely influenced by the data, therefore we considered non-informative priors for each unknown. Particularly, the log-gamma priors were chosen for the $\sigma_{\epsilon}^{-2}, \sigma_{\omega}^{-2}$ and r parametrized on the log-scale, i.e.: $\log(\sigma_{\epsilon}^{-2}), \log(\sigma_{\omega}^{-2}) \sim \log Ga(1, 5 \cdot 10^{-5})$ and $\log(r) \sim \log Ga(1, 10^2)$. Normal priors $\mathcal{N}(0, 10^3)$ were assigned for the regression coefficients and a vague normal one for the intercept.

We fitted models arising from all possible combinations of the covariates, i.e. 16384 (= 2^{14}) distinct models and ordered them according to Bayesian model comparison criteria. Particularly, the models with the best predictive performance were selected based on the lowest logarithmic score (logscore) – a measure of the predictive ability of an individual model (Ntzoufras, 2011), given by: $L_{CV} = -\sum_{s=1}^{S} \log CPO_s$, where the leave-one-out conditional predictive ordinate (CPO_s) is based on the cross-validatory predictive densities $\pi(Y_s, Y_{-s})$ for each excluded location *s*, i.e. $CPO_s = \pi(Y_s, Y_{-s})$. All the models were fitted using the R-INLA package (Rue et al., 2013) available within the R software (R Core Team, 2015). To reduce computational time, the fit of all possible combinations of covariates was done in parallel on an Intel Xeon E5-2697 CPU machine (2 × 2.60 GHz, 128 GB RAM). The fitting of 16384 models took around 2 weeks.

Subsequently, model performance was evaluated using the 5-foldcross-validation method. The dataset was randomly divided 5 times in 80% (training set) and 20% (validation set) splits of the total number of NO_2 sites. Then, models were trained on each of the five 80% subsets of the data and used to predict NO_2 concentrations at the corresponding 20% of the stations. In this case, each of the five 20% subsets is independent to the model building. The following performance metrics were examined for each fold comparing the observed NO_2 values to the posterior mean estimates (on log-scale) at the validation stations: mean absolute error (*MAE*), mean absolute prediction error (*MAPE*), root mean squared error (*RMSE*) and the coefficient of determination (R^2). The average value of the metrics over the 5 folds was presented.

We assessed the contribution of the OMI tropospheric column-tosurface conversion of NO_2 (using the surface-to-column ratios from GEOS-Chem) to the predictive ability of the GR models. In particular, two models were fitted: one with tropospheric NO_2 column densities, i.e. Ω_0 (**Model 1 - SAT**); and another one with surface NO_2 estimates based on Eq. 1, i.e. S_0 (**Model 2 - SAT-CTM**), as distinct covariates. The results were compared with the models which don't include any of the satellite-derived or CTM NO_2 proxies (**Model 0**). To evaluate whether higher spatial resolution of the CTM simulations influences the predictive ability of the models, we further fitted GR models with the CAMS-Ensemble predictor, i.e. S_{ENS} (**Model 3 - ENSEMBLE**), as an additional covariate, and compared them to the above-mentioned formulations.

Model fit and prediction was done using the SPDE method and INLA algorithm for the fast approximation of the marginal posterior distributions. In the SPDE/INLA approach, the spatial process is represented as a Gaussian Markov random field (GMRF) with mean zero and a symmetric positive definite precision matrix Q (defined as the inverse of $\Sigma_{\omega} = \sigma_{\omega}^2 R_{\omega}$). First, a GMRF representation of the Matern field was constructed on a set of non-intersecting triangles partitioning the domain of the study area (Lindgren et al., 2011). Subsequently, the INLA algorithm was used to estimate the posterior distribution of the latent Gaussian process and hyperparameters using the Laplace approximation (Rue et al., 2009). More details regarding this methodology are provided elsewhere (Blangiardo and Cameletti, 2015). Prediction was carried out after fitting the models to the full datasets (for a better spatial coverage and therefore for obtaining more accurate parameter estimates and predictions).

We used simulation-based inference to estimate the concentration of NO_2 over a gridded surface of 1 km² resolution covering the study area. In particular, 1000 samples from the posterior predictive distribution were drawn at the centroids of each grid cell (approximately 5.8 million pixels). The NO_2 maps display the sample-based medians, the standard deviations and the coefficients of variation (ratios of the standard deviation to the mean) of the posterior predictive distribution; with the last two representing a measure of uncertainty of the predictions. We determined the most polluted European areas using the first-level Nomenclature of Territorial Units for Statistics (NUTS) (EuroStat, 2019) classification of the EU to define the regions' borders. To evaluate whether the NO_2 concentration decreases as we move away from the

city centres (SimpleMaps, 2019) of each capital, buffer zones with areas varying from 1 to 30 km² surrounding the city centres were considered. The high-resolution NO_2 estimates (pixel-level posterior medians) were first clipped by each region and buffer and then averaged over the resulting sectors. For every capital, spline curves were employed to profile the relationship between pollutant concentration and distance from the city center within each buffer zone.

The Bayesian framework allowed us to make probabilistic statements about areas exceeding the international AQGs. The NO_2 concentrations exceeding the threshold limit set by the EU and WHO were calculated at pixel level by the proportion of samples drawn from the predictive posterior distributions of NO_2 that have pollution levels above the threshold. The exceedance maps were used to estimate the total number of population exposed to elevated levels of NO_2 . In particular, we overlayed the gridded population data at 1 km² spatial resolution with the threshold map and summed up all the pixels within a particular country that have exceedance probability higher than 50%. Repeating it for each of the 1000 samples drawn from the posterior distribution allowed us to estimate the number of exposed population per country together with the prediction uncertainty.

3. Results

3.1. Evaluation of GEOS-Chem and OMI tropospheric NO₂ columns

Figs. 2(a)-(b) depict yearly-averaged OMI-observed tropospheric NO2 columns and GEOS-Chem nested-grid model simulations in 2016 over the study area. In order to compare the agreement between tropospheric columns for these annual averages, the OMI observations, available at $(0.25^{\circ} \times 0.25^{\circ})$ spatial resolution, were aggregated to the extent of the nested $(0.5^{\circ} \times 0.625^{\circ})$ GEOS-Chem grid. Additionally, since the OMI observations include missing data for different days within the year (due to clouds and large solar zenith angles) whereas GEOS-Chem simulations have full daily coverage, for the computation of yearly-averaged values, the GEOS-Chem estimates were extracted only at those times that OMI data is available. Table 2 provides summary statistics for the resulting estimates (over land). There is a good agreement between the products in terms of mean, median, as well as 25% and 75% quantiles. However, the densities of the tropospheric columns obtained from GEOS-Chem are spreading over a wider range of values, as indicated by a larger standard deviation as well as lower minimum and higher maximum. Comparison of the complete GEOS-Chem dataset to the one sampled only at the locations and days when OMI is available shows that the maximum N available is the same, meaning that there is at least one OMI observation within the year for every pixel. However, the mean and the median of the full dataset is slightly higher in case of full coverage.

The spatial patterns of the GEOS-Chem and OMI columnar abundances for the study domain (i.e. over land) are highly consistent (correlation R = 0.88, N = 2193), however the tropospheric columns simulated in GEOS-Chem tend to be higher than OMI-observations over urban areas, especially over big European cities (Fig. 2(c)). The annual mean ratios of NO_2 integrated over the lowest layer of GEOS-Chem to the tropospheric column, depicted in Fig. 2(d), suggest that there is a variation in space of the vertical distribution of NO_2 , when averaged over the entire year; the values across Europe vary between 0.04 and 0.33.

3.2. Surface NO₂ retrieval

Fig. 4(a) shows the yearly-averaged OMI-derived surface NO_2 concentrations obtained using Eq. 1 (S_0) at the spatial resolution of the nested-grid simulations ($0.5^\circ \times 0.625^\circ$). The spatial pattern of the resulting estimate is similar to that observed in the original tropospheric OMI column (Ω_0) (Fig. 2(a)). However, the differences for some regions (e.g. Greater London, Madrid, Be-Ne-Lux region) are also evident. The



Fig. 2. Comparison of yearly averaged tropospheric NO_2 columns (molecules/cm² [×10¹⁵]) from OMI and GEOS-Chem (GC) in 2016. **a**: OMI at original resolution. **b**: Nested-grid GEOS-Chem model simulations sampled only at the locations and days when OMI data is available **c**: Difference between GEOS-Chem simulations sampled only at the locations and OMI resampled to the spatial resolution of the GEOS-Chem nested-grid. **d**: The ratios of NO_2 integrated over the lowest layer of the GEOS-Chem to the tropospheric column.

Summary statistics for tropospheric NO_2 annual averages for 2016 derived from OMI, GEOS-Chem nested grid simulations (in 10^{15} molecules/cm²) extracted only at the locations and days when OMI is available and GEOS-Chem simulations with full spatio-temporal coverage (i.e. with no missing).

Product	Ν	Mean	Median	SD	Min	$q_{0.25}$	$q_{0.75}$	Max
OMI NO_2 (Ω_O) GEOS-Chem NO_2 (Ω_G) GEOS-Chem NO_2 (with no missing)	2193 2193 2193	1.95 2.07 2.44	1.60 1.60 1.90	1.21 1.81 1.89	0.46 0.19 0.41	1.19 0.90 1.21	2.33 2.64 3.03	9.18 16.62 16.32

surface simulations derived from GEOS-Chem and Ensemble models, averaged over the year 2016, are illustrated in Fig. S1 (in the appendix).

3.3. Geostatistical model selection

The variable selection process for the estimation of surface NO_2 concentrations revealed that not all of the tested covariates are needed to achieve optimal predictions. The five best selected combinations of covariates (based on the lowest logscore values) for Model 0, Model 1 (SAT), Model 2 (SAT-CTM) and Model 3 (ENSEMBLE) are shown in Table 3. The model selection indicated that for both of the SAT and SAT-CTM cases, the columnar proxies derived from OMI (i.e. Ω_0) and the surface NO_2 mixing ratios (i.e. S_0) are among the covariates included in the best models, with SAT-CTM slightly outperforming the SAT formulations in terms of logscore. We also see an improvement in logscore values compared to Model 0, in which neither of the satellite-derived NO_2 proxies is considered as predictor. The use of the S_{ENS} covariate in the models lead to both Ω_0 and S_0 becoming non-important

predictors. The predictive ability in this case becomes even higher (as indicated by much lower logscore values of Model 3). The additional covariates included in the best five model combinations are similar for all four models (Table 3). Plotting the logscore values for every possible combination of covariates (Fig. 3) reveals that M3, outperforms the other models for any of those combinations, followed by M2, M1 and M0. The same figure shows that the range in improvement for the best model is highest between M3 and M2 and between M1 and M0.

The surface NO_2 concentrations were positively associated with both Ω_O and S_O covariates, with similar estimates (i.e. posterior medians) of the regression coefficients but slightly lower uncertainty (i.e. narrower Bayesian credible interval) for S_O and even lower uncertainty for S_{ENS} (Table 4). Additionally, we found a significant positive association of NO_2 concentration with the degree of imperviousness, nighttime light intensity, land surface temperature, road density and distance to the sea and a negative association of surface NO_2 with elevation, normalized difference vegetation index, wind speed and distance to the roads; the NO_2 concentration was higher for stations situated in areas dominated by urban structures and transport networks (i.e. over land cover category LC1) followed by industrial (LC2), agricultural (LC3) and forest (LC4) areas (Table 4).

The ENSEMBLE model provided the highest cross-validated R^2 values of 0.724, followed by the SAT-CTM ($R^2 = 0.715$), the SAT ($R^2 = 0.713$) and the one which does not include any NO_2 proxy from satellite or CTM, i.e. Model 0 ($R^2 = 0.711$). The out-of-sample MAE, MAPE and RMSE metrics of the predictive performance were also the lowest for Model 3. Very similar findings were observed when the 10-fold (instead of 5-fold) cross-validation was performed, or when the metrics were based on the posterior median (instead of mean) point estimates at the validation locations. The estimated range parameters (r) as well as the variances of the spatial process (σ_{ω}^2) were lower in

First five covariate combinations with the highest predictive ability (i.e. lowest logscore) for Model 0, Model 1 (SAT), Model 2 (SAT-CTM) and Model 3 (ENSEMBLE).

Model	Covariates	logscore
Model 0	IMP + DEM + NTL + LST + NDVI + WS + RD + DISS + DISR + LC	3.2789
	IMP + DEM + NTL + LST + NDVI + WS + RD + DISR + LC	3.2789
	TCD + IMP + DEM + NTL + LST + NDVI + WS + RD + DISS + DISR + LC	3.2791
	IMP + DEM + NTL + LST + NDVI + PREC + WS + RD + DISS + DISR + LC	3.2792
	TCD + IMP + DEM + NTL + LST + NDVI + WS + RD + DISR + LC	3.2792
Model 1 (SAT)	$\Omega_0 + IMP + DEM + NTL + LST + NDVI + WS + RD + DISS + DISR + LC$	3.2758
	$\Omega_O + IMP + DEM + NTL + LST + NDVI + WS + RD + DISR + LC$	3.2759
	$\Omega_O + TCD + IMP + DEM + NTL + LST + NDVI + WS + RD + DISS + DISR + LC$	3.2760
	$\Omega_0 + TCD + IMP + DEM + NTL + LST + NDVI + WS + RD + DISR + LC$	3.2760
	$\Omega_0 + IMP + DEM + NTL + LST + NDVI + PREC + WS + RD + DISS + DISR + LC$	3.2761
Model 2 (SAT-CTM)	$S_0 + IMP + DEM + NTL + LST + NDVI + WS + RD + DISS + DISR + LC$	3.2752
	$S_0 + IMP + DEM + NTL + LST + NDVI + WS + RD + DISR + LC$	3.2753
	$S_0 + TCD + IMP + DEM + NTL + LST + NDVI + WS + RD + DISS + DISR + LC$	3.2754
	$S_0 + TCD + IMP + DEM + NTL + LST + NDVI + WS + RD + DISR + LC$	3.2754
	$S_0 + IMP + DEM + NTL + LST + NDVI + PREC + WS + RD + DISS + DISR + LC$	3.2755
Model 3 (ENSEMBLE)	$S_{ENS} + IMP + DEM + NTL + NDVI + WS + RD + DISR + LC$	3.2618
	$S_{ENS} + IMP + DEM + NTL + NDVI + WS + RD + DISS + DISR + LC$	3.2619
	$S_{ENS} + TCD + IMP + DEM + NTL + NDVI + WS + RD + DISR + LC$	3.2621
	$S_{ENS} + IMP + DEM + NTL + NDVI + WS + RD + PREC + DISR + LC$	3.2621
	$S_{ENS} + IMP + DEM + NTL + LST + NDVI + WS + RD + DISR + LC$	3.2621

IMP - Imperviousness; DEM - Digital elevation model; NTL - Night time lights; LST - Land surface temperature day & night; NDVI - Normalized difference vegetation index;

RD - Road density; DISS - Distance to sea; DISR - Distance to roads; LC - land cover; TCD - Tree cover density; PREC - Precipitation.



Fig. 3. Model selection. Predictive performance of the M0, M1, M2 and M3 models (ordered according to the logscore values) arising from all possible combinations of covariates for each formulation. All the 16384 models (left) and zoom in on the best 1000 models (right).

Models 1 and 2 when compared to Model 0, indicating that inclusion of either Ω_O or S_O covariate decreases the spatial variability in the residuals. However, this variability was higher in Model 3, as shown by the higher r and σ_{ω}^2 estimates compared to all other models. This is due to the fact that the number of covariates resulting in optimal predictions is lower for Model 3; thus, the *LST* and *DISS* covariates become non-important (Table 4) in this case. The separate contributions of the spatial process (ω_i) as well as of OMI (Ω_O), OMI_GC (S_O) and Ensemble (S_{ENS}) covariates to the predictive ability of the models is presented in Table A1 in the appendix. The predictive ability using only these inputs as covariates is much lower compared to the final models that include additional climatic and land-use/cover factors.

3.4. High-resolution model-based pollutant maps and exposed population

The above results suggest that Model 3 had the best predictive ability. Therefore, statistical inferences (i.e. mapping NO_2 concentrations, exceedance probabilities and estimates of population exposure)

Posterior medians, 95% Bayesian credible intervals (BCI) and cross-validation performance metrics of Model 0, Model 1 (SAT), Model 2 (SAT-CTM) and Model 3 (ENSEMBLE) with covariate combinations giving the best predictive ability of surface NO_2 concentrations in 2016.

Covariator	Model 0	Model 1 (SAT)	Model 2 (SAT-CTM)	Model 3 (ENSEMBLE)
Covariates	Mediali (95% BCI)	Mediali (95% BCI)	Median (95% BCI)	Median (95% BCI)
Intercept	2.77 (2.62, 2.89)	2.91 (2.79, 3.01)	2.89 (2.78, 3.00)	2.80 (2.63, 2.94)
Ω_O		0.18 (0.12, 0.24)		
S_O			0.17 (0.13, 0.22)	
S _{ENS}				0.19 (0.16, 0.22)
TCD	-	-	-	-
IMP	0.11 (0.09, 0.12)	0.11 (0.09, 0.12)	0.11 (0.09, 0.12)	0.11 (0.10, 0.13)
DEM	-0.17 (-0.20, -0.14)	-0.16 (-0.19, -0.13)	-0.16 (-0.19, -0.14)	-0.16 (-0.19, -0.14)
NTL	0.21 (0.18, 0.23)	0.20 (0.18, 0.23)	0.20 (0.18, 0.23)	0.19 (0.17, 0.22)
LST	0.12 (0.07, 0.16)	0.11 (0.06, 0.15)	0.10 (0.06, 0.15)	-
NDVI	-0.07 (-0.09, -0.05)	-0.06 (-0.09, -0.04)	-0.06 (-0.09, -0.04)	-0.05 (-0.07, -0.03)
RD	0.06 (0.05, 0.08)	0.06 (0.05, 0.07)	0.06 (0.05, 0.07)	0.05 (0.04, 0.07)
SHUM	-	-	-	-
PREC	-	-	-	-
WS	-0.06 (-0.10, -0.03)	-0.06 (-0.09, -0.03)	-0.06 (-0.09, -0.03)	-0.06 (-0.09, -0.03)
DISS	0.09 (0.00, 0.18)	0.09 (0.02, 0.16)	0.09 (0.02, 0.16)	-
DISR	-0.04 (-0.06, -0.02)	-0.04 (-0.06, -0.02)	-0.04 (-0.06, -0.02)	-0.03 (-0.05, -0.02)
LC:				
LC2	-0.11 (-0.15, -0.07)	-0.11 (-0.15, -0.07)	-0.11 (-0.15, -0.07)	-0.11 (-0.15, -0.07)
LC3	-0.15 (-0.21, -0.09)	-0.16(-0.22, -0.10)	-0.16 (-0.22, -0.10)	-0.16 (-0.22, -0.10)
LC4	-0.34 (-0.42, -0.26)	-0.36 (-0.44, -0.28)	-0.36 (-0.44, -0.28)	-0.38 (-0.46, -0.30)
(1) σ_{ϵ}^2	0.10 (0.09, 0.11)	0.10 (0.09, 0.11)	0.10 (0.09, 0.11)	0.10 (0.10, 0.11)
(2) σ_w^2	0.14 (0.11, 0.20)	0.11 (0.08, 0.15)	0.11 (0.08, 0.15)	0.13 (0.08, 0.20)
⁽³⁾ r (km)	351.8 (271.4, 462.9)	286.4 (219.3, 376.5)	310.1 (236.5, 414.2)	490.9 (355.0, 684.7)
(4) MAE	0 271	0 269	0 268	0 263
(5) MAE	0.167	0.166	0.166	0.166
() MAPE	0.10/	0.100	0.100	0.100
(6) RMSE	0.363	0.361	0.360	0.356
$^{(7)} R^2$	0.711	0.713	0.715	0.724

⁽¹⁾ σ_e^2 - variance of the random error; ⁽²⁾ σ_w^2 - variance of the spatial process; ⁽³⁾ *r*- range (the distance at which the spatial variance becomes less than 10%); ⁽⁴⁾ *MAE*- Mean Absolute Error; ⁽⁵⁾ *MAPE*- Mean Absolute Prediction Error; ⁽⁶⁾ *RMSE*- Root Mean Squared Error; ⁽⁷⁾ R^2 - coefficient of determination.

are based on Model 3 with the covariate combination giving the lowest logscore (i.e. 1st among M3 - ENSEMBLE models in Table 3). Figs. 4(b)-(d) depict the predictions and their uncertainty, i.e. the median, the standard deviation and the coefficient of variation of the posterior predictive distribution for NO₂ at 1 km² spatial resolution. The most heavily polluted regions in terms of NO2 concentration include southeastern England, north-western Italy, Belgium, Netherlands, Germany (Nordrhein-Westfalen) as well as the big European cities, such as London, Rome and Paris. As one can see from the plots of the standard deviation and of the coefficient of variation, higher uncertainty is estimated at locations with higher NO2 estimates and within areas with fewer monitoring stations. Additionally, as a general trend, we found that the concentration of NO2 decreases with distance away from the centers of each European capital (Fig. S2 in the appendix). The maps of the medians, the standard deviations and the coefficients of variation of the posterior predictive distributions based on the other three models are presented in the appendix (Figs. S3-S5). Density plots of the prediction uncertainty (both, sd and cv) of the 4 models over the entire study area (i.e. over ~5.8 million pixels) show that M3 has the lowest overall uncertainty, followed by M2, M1 and M0 (Fig. 5).

Additionally, we compared estimates obtained from Model 2 with the surface NO_2 simulations from the Ensemble model. Predictions were aligned by aggregating the higher resolution estimates of Model 2 (at 1 km²) to the lower resolution of the Ensemble (at ~ 10 km²). The correlation of the predictions based on the two models was R = 0.837 at pixel level. At the monitoring locations, predictions based on the Bayesian GR model (i.e. Model 2) are closer to the measured data than the ones based on the Ensemble (Fig. 6 and Table A2). However, Model 3, that included Ensemble simulations as an additional covariate, has higher predictive ability and even lower uncertainty than Model 2.

The Bayesian framework allowed us to make probabilistic

statements about areas exceeding the international AQGs. Particularly, for NO₂ concentrations, the yearly threshold set by the European Air Quality Directive and WHO is 40 μ g/m³ (EU, 2008; WHO, 2005). This threshold is considered as an achievable objective to minimize the health impact. Fig. 7 depicts the probabilities of NO₂ concentrations exceeding the threshold for each 1 km² pixel. The vast majority of the continent meets the requirements of the AQG threshold standards, however, there are some small areas where the limits are still to be reached. The exceedance probability map (i.e. Fig. 7) was used to estimate the number of people exposed to elevated levels of NO₂ at each pixel. The estimates were aggregated at country levels and presented in Table 5 together with the prediction uncertainty (i.e. median and 95% credible intervals of the posterior predictive distribution). Results show that in 2016, 16.17 (95% C.I. 6.34-29.96) million people were exposed to NO_2 levels above the thresholds, which represent only 2.97% (95% C.I. 1.16% - 5.50%) of the total number of people living within the study area (Table 5).

4. Discussion

Nitrogen dioxide monitors do not provide full data coverage over Europe, which is an obstacle to assess the attributable health effects for the entire continent. This paper is the first to estimate surface NO_2 concentrations at 1 km² spatial resolution over 44 European countries using a hybrid approach which combines satellite-based observations, chemical transport model simulations, monitors and additional auxiliary data within a Bayesian geostatistical modeling framework.

The tropospheric columns obtained using the 3-D simulations of NO_2 concentration from the nested-grid GEOS-Chem model (at $0.5^{\circ} \times 0.625^{\circ}$ spatial resolution) were found to be highly consistent with OMI observations over the study area (correlation of R = 0.88,



Fig. 4. Surface NO_2 estimates in Europe in 2016. **a**: NO_2 mixing ratios (S_O) obtained using Eq. 1 at $0.5^\circ \times 0.625^\circ$ spatial resolution. **b**: Predicted NO_2 concentrations (i.e. median of the posterior predictive distribution) at 1 km² spatial resolution. **c**: Prediction uncertainty in terms of standard deviation (sd) of the posterior predictive distribution of NO_2 . **d**: Prediction uncertainty in terms of coefficient of variation (cv) of the posterior predictive distribution of NO_2 . Locations of the monitoring stations are overlayed on the map.



Fig. 5. *NO*₂ prediction uncertainty. Density plot of the standard deviation (left) and the coefficient of variation (right) of the posterior predictive distribution for all the 4 models at each 1 km² pixel of the prediction grid.

N = 2193). Huijnen et al. (2010) reported a spatial correlation of R = 0.8 (N = 6000) between an ensemble median of regional air quality models and Dutch OMI tropospheric NO_2 (DOMINO) v1.0.2 product (http://www.temis.nl) for the years 2008–2009, while Vinken et al.

(2014) found a higher spatial agreement (R = 0.95, N = 9270) between the nested-grid GEOS-Chem tropospheric column simulations and DOMINO v2.0 product in 2005, with both studies performed over a similar domain in Europe.



Fig. 6. Comparison between the SAT-CTM GR model estimates and the Ensemble CTM simulations. Scatter plot of the measured versus estimated NO_2 concentration at the location of the monitoring stations using the Bayesian GR SAT-CTM Model (at 1 km² resolution) and Ensemble of regional CTMs (~ 10 km²).



Fig. 7. Exceedance probabilities. Probability that NO_2 concentration exceeds the EU Directive and WHO air quality threshold of 40 μ g/m³. The zoomed areas represent some of the level 1 subdivisions of the Nomenclature of Territorial Units for Statistics (NUTS 1) classification of the European Union.

Comparison between the models M0, M1 and M2 showed that both, the satellite-derived proxies of columnar NO_2 from OMI (Ω_0) as well as the vertically corrected surface quasi-observations (S_0) obtained using surface-to-column ratios from GEOS-Chem, improved estimation of the corresponding surface NO_2 concentration and reduced prediction uncertainty, despite of missing data present in both products. This

Estimated number of people exposed to NO_2 levels above the EU and WHO thresholds in 2016 (median and 95% credible intervals of the posterior distributions).

Country	Population ¹	Exposed to NO_2
(AD) Andorra	67 462	0 (0, 0)
(AL) Albania	2 925 168	0 (0, 6136)
(AT) Austria	8731711	26 948 (3513, 71 469)
(BA) Bosnia and Herzegovina	3817952	0 (0, 0)
(BE) Belgium	11 469 758	209 841 (24 752, 535 658)
(BG) Bulgaria	7153089	0(0, 32782)
(CH) Switzerland	9 202 540	4980 (0, 41 116)
(CY) Cyprus	1 1 90 2 1 4	0 (0, 0)
(CZ) Czech Republic	10618625	2648 (0, 38496)
(DE) Germany	81 848 649	1 148 965 (418 143, 2 558 156)
(DK) Denmark	5766524	0 (0, 0)
(EE) Estonia	1 349 753	
(FL) Greece	11 050 816	99,697 (0, 1,704,795)
(ES) Spain	44 529 778	1 895 762 (581 991 3 196 598)
(ED) Finland	5 979 902	0 (0 0)
(FR) France	65 346 726	3 339 836 (885 080 5 005 125)
(CC) Guernsov	53 470	0 (0, 0)
(GI) Gibraltar	21 222	
(HP) Croatia	4 221 223	4400 (0, 108 244)
(HL) Hungary	4 221 001	4490 (0, 108244)
(HU) Hungary	1002/750	0 (0, 225 219)
(IE) Ineland	4 014 031	0 (0, 0)
(IS) Iceland	330 470	0 (0, 0)
(II) Italy	60 499 999	4 620 219 (2 964 773,
	00 550	6 229 796)
(JE) Jersey	92 559	0 (0, 0)
(LI) Liechtenstein	3/ 303	0 (0, 0)
	2 923 123	0 (0, 0)
(LU) Luxembourg	5/0 985	0 (0, 0)
(LV) Latvia	2 125 289	0 (0, 0)
(MC) Monaco	13 681	0 (0, 0)
(ME) Montenegro	664 404	0 (0, 0)
(MK) North Macedonia	2123204	0 (0, 3780)
(MT) Malta	420 872	0 (0, 0)
(NL) Netherlands	17718000	140 468 (25 726, 527 690)
(NO) Norway	5 372 166	6015 (575, 15649)
(PL) Poland	39 1 49 578	70 633 (6204, 457 799)
(PT) Portugal	9 700 000	0 (0, 13646)
(RO) Romania	19740811	196 024 (0, 1 253 267)
(RS) Serbia	8 995 232	96 610 (0, 475 274)
(SE) Sweden	10 521 396	0 (0, 0)
(SI) Slovenia	2112593	0 (0, 17 407)
(SK) Slovakia	5 486 419	0 (0, 3972)
(SM) San Marino	30 187	0 (0, 0)
(UK) United Kingdom	65 644 463	4 304 920 (1 422 934,
		7 431 681)
(VA) Vatican	1970	1970 (1970, 1970)
Whole study area (44 countries)	544 472 605	16 170 026 (6 335 661,
		29 955 725)
		2.97% (1.16%, 5.50%)

 $^1\rm Estimate$ obtained via cubic spline interpolation of 2000, 2005, 2010, 2015 and 2020 population data at 1 $\rm km^2$ pixel level.

important contribution of satellite-derived NO_2 proxies to the estimation of surface concentrations was also found using LUR models in Europe (Vienneau et al., 2013), US (Novotny et al., 2011) and globally (Larkin et al., 2017). Although the above-mentioned works have not assessed the contribution of the vertical correction itself to the model predictive ability, several recent works (Bechle et al., 2013; Knibbs et al., 2014; Bechle et al., 2015), suggested that the satellite column abundances from OMI (Ω_0) may be sufficient to estimate the spatial patterns of ground-level NO_2 , therefore conversion of columnar values to surface concentrations becomes unnecessary within a LUR model. In fact, Knibbs et al. (2014) found that the annual NO_2 models that included NO_2 tropospheric column as covariate exhibited a slightly better predictive ability compared to those that included estimates of surface NO_2 obtained by modelling surface-to-column ratios using the Weather Research and Forecasting (WRF-Chem) chemical transport model for a LUR study in Australia. Bechle et al. (2015) reported similar findings using DOMINO tropospheric columns and GEOS-Chem surface-to-column ratios for a LUR study in the US. Our results showed that this conversion improved the predictive ability of the geostatistical models. However, the improvement was marginal when additional predictors were considered in the models, as indicated by the logscore values (differences in the 4th digit) and the cross-validation performance metrics (i.e. R^2 , *MAE* and *RMSE* presented here on the log-scale).

To further evaluate, whether the vertical correction is more important at locations with large variability in the surface-to-column ratio (Fig. 2(d)), we extracted the ratio values at the location of the stations and re-fitted the models M0–M2 on the subset of the data with (i) large variability (i.e. ratio smaller than 0.25th quantile or higher than 0.75th quantile of all the ratio values) and (ii) almost constant values (i.e. ratio higher than 0.25th and lower than 0.75th quantile of all the values). The results (Table A3 in the appendix) revealed that, indeed, the vertical correction using CTM simulations results in a higher predictive performance when the variation in the ratio is high. Therefore, conversion of the columnar values to surface concentrations is more important in cases when a particular spatio-temporal domain of interest ascertains higher variation in the ratio (e.g. in case of daily or weekly, rather than annually averaged data).

Rigorous variable selection indicated that additional high-resolution auxiliary earth observation data such as the novel pan-European Copernicus land monitoring products, including the impervious surfaces and land-use/cover datasets were important predictors in all models. The road density and distance to the roads covariates computed from the OpenStreetMap project as well as the distance to the sea, elevation, night time lights intensity, NDVI, LST and wind speed increased the predictive ability of the best model (based on the logscore). The positive/negative associations of these covariates with surface NO₂ measurements from monitoring stations generally agree with those reported in the literature. In particular, the important positive associations of surface NO2 with impervious surfaces and major roads were also found in Novotny et al. (2011), Knibbs et al. (2014), Bechle et al. (2015), while the negative effect of NDVI was also estimated in Larkin et al. (2017). However, in contrast to studies in the US (Novotny et al., 2011; Bechle et al., 2015), we found a positive effect of the distance to the sea (or coast) on NO2 and a negative association with elevation, which is consistent with another European study by Beelen et al. (2009).

The comparison of the simulations from the Ensemble of regional CTMs (i.e. S_{ENS}), available at 10 km² resolution, with the estimates obtained using Model 2 (SAT-CTM) aggregated at the same scale has shown a good spatial agreement. Extracted at the locations of the stations, the 1 km² predictions from Model 2 were closer to the NO_2 measurements than the S_{ENS} simulations, implying that the higher the resolution of the estimated NO_2 concentration is (i.e. of the gridded GR predictions or CTM simulations), the smaller the underestimation of the exposure gets. When S_{ENS} was included as a covariate in the GR model (i.e. Model 3), its predictive ability increased and the prediction uncertainty reduced, granting the most accurate surface NO_2 estimates.

The advantage of the geostatistical models is their ability to provide information in areas where there is no monitoring (i.e. no stations). Traditional statistical approaches that are used to produce gridded estimates (e.g. LUR or GWR) usually fail to quantify the uncertainties in the predictions (which are necessary for studies related to health risks assessments). For a single CTM, sensitivity analyses (Saltelli et al., 2004) quantify uncertainty by e.g. perturbing the emissions and meteorological conditions. For multi-model ensembles of CTMs (Galmarini et al., 2018; Marécal et al., 2015; Potempski et al., 2008), the analyses from different models (varying in number) provide a composite of model simulations (e.g. medians or percentiles). In this case, the final estimates at the various points in space and time are more conservative, filter the single model differences and are usually more consistent with the actual evolution of the process (Galmarini et al., 2004). On the other hand, Bayesian inference treats the unknown NO_2 concentrations at pixel level as random quantities and estimates their predictive posterior distribution, a probability distribution that incorporates the uncertainty of the predictions via its variance. Sensitivity analysis in the context of Bayesian modelling assesses influence the alternative distributions (for the data or the priors) have on the predictions. By drawing samples from the predictive posterior distribution, we can estimate the mean, median, standard deviation or any other quantity of interest (e.g. probability of exceeding any particular threshold). The assessment of exposure burden also becomes possible by utilizing gridded population data.

Estimates from the posterior predictive distribution show that the most heavily polluted regions in terms of NO2 are south-eastern England, north-western Italy, Belgium, Netherlands, Germany (Nordrhein-Westfalen) as well as the big European cities, like London, Brussels, Hamburg, Berlin, Paris and Athens. Similarly to Shaddick et al. (2013), we found that higher uncertainty of the predictions is estimated in areas with elevated pollutant concentrations (as seen in the plot of the predictive standard deviation, i.e. Fig. 4(c)). A plausible explanation is that the number of stations in areas dominated by urban structures and transport networks is usually lower, which leads to higher uncertainty in the coefficient estimates of the land-cover covariates defining these regions, and therefore to higher uncertainties in areas with higher NO2 concentration estimates. Additionally, the map of the coefficient of variation (Fig. 4(d)) has shown that the predictive uncertainty is higher in regions (and countries) with fewer monitoring stations. Thus, the results of our study can help identifying areas where new monitoring stations are needed. In terms of population exposure, we found that only about 3% of people living within the study domain were exposed to NO2 values exceeding the international AQGs. This result suggest that NO₂ represents a much smaller threat to Europe, when compared to particulate matter (both fine - PM2.5 and coarse - PM_{10}) concentrations. For the latter, it was estimated that more than 66% of people within the same area and year were breathing air abvove the WHO thresholds (Beloconi et al., 2018). Our results are comparable to findings put forth in a recent EEA Air Quality report (European Environment Agency, 2019) based on Eionet gridded estimates (using a completely different approach) for 2016 (Eionet, 2018). Indeed, it was reported that 2.8% of population in almost the same area covered by our study were exposed to NO2 concentrations above the WHO/EU thresholds.

The EU Directive sets criteria for the minimum number of the sampling points and for the site locations. They require Member States to locate sampling points both "where the highest concentrations occur" (with traffic or industrial type stations) and in areas which are "representative of the general population's exposure" (with background type stations). However, the selection of site locations involves a degree of flexibility, which allows Member States to not necessarily measure air quality near major industries or main urban traffic. Additionally, the Air Quality Directive requires that Member States maintain sampling points that have exceeded PM_{10} , but this obligation does not apply to other pollutants, in particular to NO_2 (European Court of Auditors, 2018). Statistical models rely on the monitoring data, and therefore they may underestimate the concentration of pollutants, particularly in urban areas.

Our methodology can be applied to estimate NO_2 concentration and evaluate international AQGs for any specific year or spatial domain. The operational application of the models for a different year or area of investigation although possible, is not straightforward, since it requires re-estimation of the regression parameters and spatial process, based on the availability of the station data and the covariates. The dense *in situ* network available in Europe allows for an accurate estimation of the spatial correlation structure, especially over shorter distances. In other world regions, except North America and Asia, the station coverage is much lower (Larkin et al., 2017). Reliable estimation of the spatial correlation is not possible due to the large distances between the stations. Therefore, we expect that the satellite/CTM covariates will improve model predictions even more than in Europe.

The number of days with missing OMI varies in space. In particular, higher proportion of missing values is observed in the northern part of the continent and at high altitudes due to the presence of clouds, higher surface reflectance and larger solar zenith angles. Our results have shown, that even with missing values, the annual averaged OMI product was a significant predictor either vertically corrected (Model 2) or not (Model 1). In order to evaluate the impact of missing OMI data on the predictive ability of Models 1 and 2, we re-fitted the models to the subset of stations located in pixels with larger and with lower availability of OMI. We defined the availability of OMI as the proportion of days with complete data and considered that this is "large" when the proportion is above the median of those across the stations (here 39%). The results (Table A4 in the appendix) have shown that the contribution of OMI covariate to the predictive ability of the model was higher when models were fitted to data from stations with large availability (i.e. proportion of days with complete data >39%). Therefore, the number of missing data influences the predictive ability of the models. Methods for missing satellite-data imputation could improve estimates. The frequentist statistical/machine learning methods (e.g. random forests) available for filling the data gaps cannot quantify the uncertainty in the estimates, which propagates to the NO2 outcome and therefore complicates results interpretation. Bayesian models can estimate both the missing values and the model parameters within a single hierarchical formulation. However, this is computational not yet feasible for very large surfaces such as our study region. Furthermore, whereas the annual EEA station data were calculated based on the daily averaged values, the daily OMI measurements and GEOS-Chem simulations are confined to the OMI overpass window (~ 13-15 LT). This may also explain the small contribution of OMI NO₂ covariate to the predictive ability of the models, given the diurnal variation (cycle) in NO₂ concentrations. For specific applications, such as those health related, it is important to estimate the total annual average, rather than NO₂ average at a particular time within the day. Another drawback is the varying spatial resolution of the covariates. We analysed data at the original georgraphical scales assuming homogeneity of their values within a coarser spatial resolution. Approaches have been proposed to address the fusion of data with different spatial supports (Nguyen et al., 2012; Beloconi et al., 2016) and particularly for statistical downscaling of coarse resolution CTM outputs (Fuentes and Raftery, 2005; Berrocal et al., 2010).

While the OMI's 15-year dataset (2004-onwards) is suitable for investigating past NO_2 changes and trends, the high spatial resolution of the recently released Sentinel 5P/TROPOMI (7 × 3.5 km² near nadir) (Veefkind et al., 2012) will allow detection of small-scale NO_2 sources and will increase the fraction of cloud-free observations by an estimated 70% as compared to OMI (Krijger et al., 2007). The results presented here are very promising, especially in view of these new data. As observed in this work, higher resolution CTMs lead to better estimates, therefore we expect that the Sentinel 5P/TROPOMI measurements, either alone, or combined e.g. with the CAMS-Ensemble model will lead to even higher predictive ability.

5. Conclusions

We have shown the benefits of combining data from monitors, satellites and chemical transport models in a rigorous Bayesian geostatistical framework to estimate the spatial distribution of NO_2 in Europe. We presented the results based on different input data (satellite, global and regional CTMs) and evaluated the contribution of the column-tosurface conversion of satellite products and of the effect of spatial resolution of CTMs to the predictive ability of the GR models. The results indicated that the vertical correction slightly improves the predicted estimates, whereas higher resolution CTMs improve the accuracy of the prediction. Based on the best model and available monitoring data, we evaluated the AQGs and estimated the number of people living in areas exceeding the thresholds. Due to data limitations, such as coarse number of stations in some areas and the complexity of NO_2 variability in urban environments, the findings might underestimate the population exposure, particularly in urban areas. However, our work is the first to estimate the population exposure to NO_2 in 44 European countries at 1 km² resolution based on Bayesian geostatistical modelling, which explicitly quantifies the prediction uncertainty.

CRediT authorship contribution statement

Anton Beloconi: Conceptualization, Methodology, Software, Validation, Formal analysis, Investigation, Data curation, Writing original draft, Writing - review & editing, Visualization. **Penelope Vounatsou:** Conceptualization, Methodology, Resources, Data curation, Writing - review & editing, Supervision, Project administration, Funding acquisition.

Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Appendix A. Supplementary material

Supplementary data associated with this article can be found, in the online version, at https://doi.org/10.1016/j.envint.2020.105578.

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